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RESEARCH ARTICLE: Genesis, classification and evaluation of some sugarcane growing black soils in semi arid tropical region of Telangana

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BACKGROUND AND OBJECTIVES

Land is a finite natural resource and there is little scope to increase the areas under cultivation. The efforts made in the past to bring new areas under cultivation which results the reduction forests to 20% of total geographical area of the country. For decades, advancement in agricultural practices has been a necessity due to ever increasing demand caused by growing population. Sugarcane (*Saccharum officinarum* L) is

being cultivated is India is an area of 42.45 lakh ha in the states of Karnataka, Maharashtra, Madhya Pradesh, Telangana, Andhra Pradesh, Tamil Nadu, Goa and Kerala with total sugar production of 192.67 lakh tones. India is one of the largest producers of sugar and shares about 41.11 % and 13.25 % of Asian and Worlds sugar production, respectively. The population in India is increasing steadily and as well the demand for sugar and other sweetening agents because of changing food habits. There is no scope to increase the area under sugarcane to meet the requirements. This envisages the adoption of better crop production and protection technologies for increased production per unit area and time apart from varietal improvement. Sugarcane is cultivated in the Medak district in an area of 22076 hectares producing 1721 thousand tonnes with an average productivity of 74.41 t ha⁻¹.

Telangana state being under a semi-arid tropical monsoon climate, has a number of soil types which are found in all types of climates, occupying 3.5 per cent (114,840 sq km or 114.84 lakh ha or 11.484 m.ha) of the country's geographical area. Hence their management varies from place to place besides the crop variation. Maintaining the soil with high productivity on sustainable basis is important to meet basic needs of the people. Hence delineating the sugarcane growing soils for their fertility helps in understanding the soil related constraints and their intensity which is essential to develop site specific management strategies.

Classification of sugarcane growing soils in a taxonomic perspective provides information on the nature and its potential production capabilities. The characterization and classification of soils helps in determining the soil potential, identifying constraints and giving detailed information on different soil properties of the sugarcane growing areas. The present investigation is aimed to assess the characteristics of soils and land resources to comprehend the potential capability of sugarcane growing soils of Medak district in the perspective of developed land use decision for effective utilization of resources.

RESOURCES AND **M**ETHODS

Location and brief description of the study area :

Medak district of Telangana state is extending over an area of 9,519 km². It forms a part of Deccan Plateau under Godavari basin and lies between North Latitudes 17^o 27' and 18^o 18' and East longitudes 77^o 28' and 79^o 10' falling in topographical sheet nos. 56 F, G, J and K of Survey of India. It is bounded by the Nizamabad district on the north, Karimnagar district on the north and northeast, Warangal and Nalgonda district on the east, Hyderabad and Rangareddy district on the south and Bidar district (Karnataka) on the west. It is divided into three revenue divisions, *viz.*, Sangareddy, Medak and Siddipet with 46 revenue mandals/tehsil and 1223 villages in the district. The district is divided into 12 agricultural divisions.

Based on the morphological characteristics and physiography, four geo-referenced pedon were selected in Medak district such as Aroor (Pedon 1), Budera (Pedon 2), Andole (Pedon 3) and Pulakurty (Pedon 4). Physiographically, the district forms part of South Deccan Plateau. It is an ancient plateau exposed for long ages to denudation. Sheet-wash and retreat of hill slopes are the major geomorphic processes responsible for sculpturing of the present day landforms under semi-arid conditions. The plateau has two erosional surfaces with altitudes of 150-600 m and 300-900 m above MSL. It has been divided into three physiographic regions, viz., granite and granite-gneiss landform, basalt landform and laterite landform. The important rock types are Peninsular Gneissic complex, Dharwar super group associated with Younger intrusive of Achaean age separated unconformably with overlying Basaltic flows of late Cretaceous to early Eocene age with sub-Recent to Recent alluvium along the stream courses. The Archaean or Peninsular gneisses occur all over the district in 6, 86,853 ha area (70.7%). They are partially metamorphosed igneous rocks. They remained stable as a "Shield" area for a very long time. The rocks are composed of grey or pink feldspars, quartz and muscovite mica (NBSS&LUP, 2005).

The climate is semi-arid. The mean annual rainfall is 870 mm of which 76 per cent is received during the southwest monsoon (June to September), 14 per cent during the northeast monsoon (October to December) and 8 per cent during the premonsoon period (March to May). The rainfall is highest in the month of August. The climate of Medak district is comparatively equitable and although it is very hot in May with mercury rising up to 42° C. The temperature dips to 12°C in winters during the months of December and January. The mean maximum and minimum temperature vary from 40° to 26°C. Mean humidity varies from 65 per cent in July to 74 per cent in December. The soil moisture content is dry for more than 90 cumulative days or 45 consecutive days in the months of summer solstice. The soil moisture and temperature regimes of the study area are Ustic and Isohyperthermic, respectively. The natural vegetation existing in the study area are grasses, shrubs, thorny bushes such as *Cynodon dactylon, Cyprus rotundus*, *Butea frondosa, Dalbergia latifolia, Azadirachta indica, Tectona grandis, Terminalia tomertose and Acacia spp. Prosopis juliflora, Cacia sp*, broad leaf weeds such as *Selotia, Parthenium, Eucalyptus, Euforbia* sps., *etc.* The principal crops cultivated are Rice, Maize, Sugarcane, cotton, redgram, Greengram, Blackgram, Groundnut and potato.

Collection and processing of soil samples :

The geo-referenced black soil pedons were selected on the basis of soil heterogeneity and land forms in different locations of sugarcane growing areas of the district. The profiles were dug up to parent rock and studied for their morphological characteristics (Soil Survey Division Staff, 1993). Soil samples were collected from each horizon of four representative pedons for laboratory analysis. The samples were air dried and ground to pass through 2 mm sieve. Relevant physical and chemical properties were determined by following standard analytical procedures. The Soil pH, EC (1:2.5 soil water suspension); exchangeable cations (Jackson1973); cation exchange capacity (Chapman, 1965); organic carbon (Walkly and Black, 1934); free CaCO3 (Piper 1966); bulk density (Blake and Hartze 1986); moisture retention at 33 and 1500 kPa (Richards, 1954); COLE (Soil Survey Staff, 2010); water holding capacity and volume expansion (Sankaram, 1966); gravel by gravimetry method (Govindarajan and Koppar, 1975); Soil texture by International Pipette Method (Piper, 1966). The soils were characterized and classified as per Keys to Soil Taxonomy (Soil Survey Staff, 2010).

OBSERVATIONS AND ANALYSIS

The results obtained from the present study as well as discussions have been summarized under following heads :

Morphological properties :

Major part of the study area is coming under nearly

level to gently sloping. The details of the morphological properties of the soils were presented in Table 1. The drainage condition of the pedon 1, 2, 3 and 4 was poorly drained with irregular CaCO₂ concretions. Soil depth of the pedons 1, 3 and 4 were very deep (>150 cm) and deep soils were found in nearly level to very gently sloping plain. The pedon 2 was moderately deep (75-100 cm), The variation of depth in relation to physiography, mainly because of non-availability of adequate amount of water for prolonged period on upland soils associated with removal of finer particles and their deposition at lower pediplain have resulted in shallow soils in uplands and deeper soils in lowland physiographic units. The results obtained in the present study are in agreement with the findings of Ramprakash and Seshagiri Rao (2002). All the pedons were characterized as A- B-C horizons. The thickness of the surface horizons varied from 14.0 (pedon 2) to 27.0cm (pedon 4) cm and sub surface horizons ranged from 17.0 (pedon 2) to 40.0 cm (pedon 3). Pedon 1, 3 and 4 showed slicken sides (Bss). The surface horizon was designated as 'Ap' horizon at all the location because of the ploughed and disturbed condition due to cultivation. Similar observation was made by Rajeshwar et al (2009) and Ashok kumar and Jagadish Prasad (2010) to represent ploughed condition of the soils.

The 'B' horizon of soils (pedon 1, 9 and 12) exhibited prominent, well-formed distinct slickensides. Hence the symbol 'ss' (sub-ordinate distinction) was suffixed to the master horizon symbol 'B'. Similar type of designation was represented by Chinchmalatpure et al. (2005); Ashok kumar and Jagadish Prasad (2010); Rajeshwar and Mani (2015). The boundary between the sub-horizons of 'Bss' horizon was described as diffuse because of the presence slickensides and the clay content was high enough for clay textural class. Similar finding was noticed by Balapande et al. (2007). The soil colour varying from very dark grayish brown (10 YR3/2) to dark yellowish brown (10YR 4/4) under dry condition and very dark gray (10 YR3/1) to dark yellowish brown (10YR 4/4) under moist condition. The dark colour appears due to the presence of iron and manganese oxide in combination with the organic complex (Srinivasan et al., 1969) and parent material, topography, high clay content, clay-humus complex, smectite type of clay, moisture etc., were the factors responsible for the dark colour of the soils (Santsingh, 1987).

The surface horizons and subsurface horizons had

Table 1	Table 1: Morphological characteristics of Sugarcane growing soil pedons of the Medak district	ch ar acteri st	tics of Suga	rcane growi	ng soil pe do	ns of the M	ledak distric	t								
Pedon	Location	Horizon	Depth (cm)	C ₀ Dry	Colour Moist	- Texture	Structure	Dry	Consistency Moist V	ncy Wet	Efferv es cence	Pores	Roots	Boundary	Cutans	Other features
	Sadasivpet division	sion														
1.	Aroor	Very-Fin	te Smectitic	Calcareous S	Superactive I	sohy perther	Very-Fine Smeetitic Calcarcous Superactive Isohyperthermic Typic Haplusterts (Latitude ^o N 17 ^o 37'42.34" Longitude ^o E 77 ^o 53'10.33" Attitude 539 m)	apluste	arts (Lat it	ude [°] N 17	1° 37'42.3	4" Long	tude [°] E	77° 53' 10.33	"Altitude	539 m)
		Ap	0-25	10YR3/2	10YR3/2	c	m labk	sh	ų	vs&vp	ms	ff	mf	cs		Gilgai relief;
		BA	25-52	10YR3/2	10YR3/1	c	m2abk	h	vfi	vs&vp	ms	ff	cf	Sô		surface cracks
		Bss1	52-79	10YR3/2	10YR3/2	c	m2abk	h	vfi	vs&vp	ms	ff	ı	cd		(5-8 cm wide);
		Bss2	79-115	10YR4/2	10YR3/2	c	m2abk	Ч	vfi	vs&vp	s	ff	ı	cd		slickensides;
		Bss3	115-155	10YR4/2	10YR3/2	c	m2abk	ų	ų	vs&vp	VS	ff		cd		sub surface hard
		С	155 +	Mixed with	Mixed with calcareous murram	nurram										pan with high
		;						,	ļ							B.D
2.	Budera	Fine Sm	ectitic Super	active Non c	alcarious Iso	hypertherm	Fine Smeetic Superactive Non calcarious Isohyperthermic Vertic Haplustepts(Latitude "N 17" 38'37.18" Longitude 77" 50'35.00" Altitude 585 m)	plustep	ts(Latitu	lde °N 17	^o 38'37.18	"Longit	ude ^{°77°}	50'35.00"AJ	titude 585	m)
		Ap	0-14	10YR3/2	10YR3/1	cl	f1 sbk	sh	ſŕ	s&p	ms	ff	mf	cs		subsurface
		Bwl	14-41	10YR3/2	10YR3/1	cl	m1sbk	Ч	fr	vs&p	ms	ff	ff	сw	,	cracks(3.4cm)
		Bwss1	41-58	10YR4/2	10YR4/1	cl	m 1 sbk	Ч	ĥ	vs&p	ms	ff	ff	cw	,	& slicken sides
		Bwss2	58-79	10YR4/2	10YR4/2	cl	m 1 sbk	Ч	ĥ	vs&p	ms	ff	ı	CW		
		BC	79-100	10YR4/4	10YR4/4	cl	m1sbk	Ч	ĥ	s&p	ms	ff		cw		
	Jogipet division	_														
3.	Andole	Fine Sme	xtiti Calcan	ous Superac	tive Isohyper	thermic Ty	Fine Smeet it Calcareous Superactive Isohyperthermic Typic Haplusterts (Latitude ^o 17 ^o 49 ³ 34.54 ⁿ Longitude ^o 78 ^o 05 ^o 7.31 ⁿ Altitude 492 m)	tts(Lat	itude [°] 1'	7°49'34.5	4" Longit	ude [°] 78 [°]	05'07.3	l" Altitude 49	92 m)	
		Ap	0-25	10YR3/2	10YR3/2	c	m2sbk	h	ĥ	vs&vp	ms	ff	mf	cs		Gilgai relief
		\mathbf{BA}	25-65	10YR3/2	10YR3/2	c	c2abk	h	ĥ	vs&vp	ms	ff	ff	cs		surface cracks
		Bss1	55-85	10YR3/2	10YR3/2	c	c2abk	Ч	vfi	vs&vp	ms	ff	ff	cd	,	(5-8 cm wide)
		Bss2	85-117	10YR3/2	10YR3/2	c	c2abk	Ч	vfi	vs&vp	SS	ff	ı	cd	,	and subsurface
		Bss3	117-145	10YR3/2	10YR3/2	c	c2abk	h	vfi	vs&vp	SS	ff	ı	cd	1	cracks(3-5cm)
		Bss4	145-178	10YR3/2	10YR3/2	c	c2abk	Ч	vfi	vs&vp	SS	ff	ı	cd		slickensides
																conca
	Narayankhed di visi on	li visi on														
4.	Pulkurty	Very- Fi	ne Smectitic	, Calcareous	Superactive	Isohypert h	Very-Fine Smectric, Calcareous Superactive Isohypethemic Typic Haplusterts (Latitude ° 17° 56' 45.77" Longitude 77° 42' 43.82" Altitude 527 m)	Haplus	terts(Lat	itude°17	° 56'45.77	" Longit	ude° 77°.	42'43.82" AJ	titude 527	m)
		Ap	0-27	10YR3/2	10YR3/2	c	m2sbk	Ч	ĥ	vs&vp		ff	cf	cs		Gilgai relief
		\mathbf{BA}	27-55	10YR3/2	10YR3/2	o	c2abk	Ч	ų	vs&vp	sm	ff	cf	cs		surface cracks
		Bss1	55-87	10YR3/2	10YR3/2	c	c2abk	Ч	ĥ	vs&vp	ms	ff	ff	cs		(5-7 cm wide);
		Bss2	88-124	10YR3/2	10YR3/2	c	c2abk	Ч	ĥ	vs&vp	s	IJ	ı	cd		slickensides;
		Bss3	124-150	10YR3/2	10YR3/2	с	c2abk	h	ĥ	vs&vp	s	ff	ı	cd		sub surface hard
																pan with high B.D.
Soil texture Soil Structure Soil Consistence	ure : cture : sistence :	ls – Loamy c-coarse, m l- loose, sh-	sand, sl- Sa - medium, - slightly han	ls – Loamy sand ,sl- Sandy loam, scl- c-coarse, m- medium ,f- fine , l - wea l- loose, sh- slightly hard, h- hard ,vh-	1—Sandy clay sak, 2- moder h- very hard,	/ loam, sc- ate,3 - stro	Sandy clay loam, sc- Sandy clay, cl- Clay loam andc- Clay ç 2- moderate,3 - strong, gr- granular ,abk- angular blocky, vey hard,vfr-very friable, fr- friable , fr- frim, vf- very fim	cl- Clay lar ,abl	r loam an c- angula firm, vf-	de- Clay r blocky, s very firm	bk- sub-an , so - non	ngular bl sticky, s	ocky s-slight h	y sticky, s- st	icky,vs- ve	ls - Loamy sand, sl- Sundy loam, scl-Sandy clay loam, sc- Sandy clay, cl- Clay loam and c- Clay c-coarse, m- medium, f- fine, 1- weak, 2- moderate, 3 - strong, gr- granular, abk- angular blocky, shk- sub-angular blocky 1- bose, sh- slightly hard, h- hard, yh- very hard, yfvery firiable, fr- firm, vf- very firm, so - non sticky, ss-slightly sticky, s- sticky, vs- very sticky, po- non
Pores		Size f-fine, ps-	- sugarty pa m-medium,	suc, p-piasu c-coarse; Qu	plastic, ps – sugirity plastic, pplastic, vp- very plastic Size f-fine, m-medium, c-coarse; Quantity f-few, c-common, m-many f-few, f-f-	e-common	, m-many									
Koots Effervescence	: cence :	Nize I-Ime, m-mild ,ms	MZE I-Ime, m-medium, c-coars m-mild ,ms-moderately strong	c-coarse; Un strong s-stro	Mze I-Ine, m-medium, c-coarse; Quantity I-Iew, c-common, m-many m-mild ,ms-moderately strong ⊱strong vs-very strong	c-common trong	, m-many									
Boundary		c- clear, d-	c- clear , d- diffuse, s- smooth		,w- wavy, g- gradual, a- abrupt	ıal, a- abruț	J.									

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blocky structure (sub-angular and angular) and the peds were medium to coarse in size with strong grade (strength). The pedality of black soils was more strongly developed because of the high clay content, CEC, BSP and dominance of montmorillonite type of clay. Stronger pedality of soils at lower topographic positions might be due to finer fractions (Shyampura *et al.*, 1994). The textural class of fine earth fraction was clay loam (pedon 2) to clay (pedon 1, 3 and 4) throughout the depth. The uniformity in texture was due to the argillopedoturbation operating in the black soil profiles (Buol *et al.*, 1998); Marathe *et al.* (2003) and Balapande *et al.* (2007). The horizons of soils exhibited sticky to plastic to very sticky to very plastic, firm and hard to very hard in wet, moist and dry conditions, respectively, which might be due to high clay content. Similar observations made by Sarkar *et al.* (2001) in soils of lower outlier of Chhotanagpur plateau and Rajeshwar and Mani (2013).

Vertic properties like surface cracks ranging from 3-8 cm wide, slickensides, microknolls and microridges were developed in the soil pedons (1, 3 and 4). The soil pedons (1, 2, 3 and 4) had shown a prominent gilgai

Pedo	Locatio	Horizon	Depth	$C_{\rm max} = 1.0(1)$	Particles	size distribu	tion (%)	B.D	Pore space	W.H.C	Volume	COL
n	n	Horizon	(cm)	Gravel (%)	Sand	Silt	Clay	(Mg m-3)	(%)	(%)	expansion (%)	Е
	Sadasivj	et di visi on	L									
1	Aroor	Very-Fine	Smectitic (Calcareous Supe	eractive Iso	hypertherm	піс Туріс Н	laplusterts (La	titude ° N 17 ° 3	37'42.34" L	ongitude°E77°	
		53'10.33"	Altitude 53	39 m)								
		Ap	0-25	11.50	20.1	19.8	59.6	1.56	55.0	48.0	23.1	0.14
		BA	25-52	10.50	19.0	18.4	62.1	1.59	53.0	51.0	23.7	0.16
		Bss1	52-79	9.00	17.3	17.3	63.9	1.61	50.0	49.0	25.5	0.18
		Bss2	79-115	9.50	16.8	17.2	64.3	1.65	48.0	46.0	26.4	0.19
		Bss3	115-155	8.30	16.4	20.0	61.1	1.68	46.0	45.0	28.4	0.21
		С	155 +	Mixed	l with calca	reous murr	am					
2	Budera	Fine Smee	ctitic Supera	active Non calc	arious Isohy	perthermic	e Vertic Ha	plustepts (Lat	it ude $^{\circ}$ N 17 $^{\circ}$ 3	8'37.18" Lo	ngitude $^\circ$ 77 $^\circ$	
		50'35.00"	Altitude 58	5 m)								
		Ap	0-14	12.30	41.2	18.6	38.7	1.43	46.0	32.0	23.10	-
		Bw1	14-41	14.00	42.3	17.3	40.1	1.52	52.0	38.0	26.24	-
		Bwss1	41-58	14.50	43.2	18.6	37.2	1.50	55.0	36.0	27.13	-
		Bwss2	58-79	17.00	42.5	17.9	39.3	1.61	61.0	38.0	29.14	-
		BC	79-100	17.60	39.9	17.2	42.1	1.62	62.0	39.0	29.10	-
	Jogipet	li visi on										
3	Andole	Fine Smee	titi Calcare	ous Superactive	e Isohypertl	nermic Typ	ic Hapluste	erts (Latitude '	° 17 ° 49'34.54'	'Longitude	° 78 ° 05' 07.31" Al	t it ude
		492 m)										
		Ap	0-25	10.9	28.2	16.5	54.2	1.51	45.0	44.0	22.8	0.13
		BA	25-65	10.4	25.4	16.9	57.6	1.50	48.0	47.0	23.7	0.14
		Bss1	55-85	11.1	23.2	17.4	58.9	1.54	51.0	49.0	25.8	0.16
		Bss2	85-117	12.6	20.1	18.0	61.7	1.58	52.0	51.0	26.8	0.18
		Bss3	117-145	12.3	19.0	17.7	63.1	1.58	52.0	48.0	30.4	0.18
		Bss4	145-178	12.5	18.8	19.6	61.4	1.60	55.0	53.0	30.1	0.17
	Narayan	khed di visi	ion									
4	Pulkurt	Very-Fin	e Smectitic,	, Calcareous Suj	peractive Is	oh ypert her	mic Typic	Haplustrets (L	atitude ° 17°56	5'45.77" Loi	ngitude° 77° 42' 43	3.82"
	у	Altitude 5	27 m)									
		Ap	0-27	10.0	10.3	17.5	71.4	1.49	45.0	46.0	23.5	0.14
		BA	27-55	9.4	8.2	18.1	73.4	1.55	48.0	47.0	24.6	0.15
		Bss1	55-87	8.9	6.3	18.4	75.1	1.62	51.0	49.0	26.8	0.17
		Bss2	88-124	10.1	4.3	19.5	75.7	1.66	53.0	51.0	28.6	0.19
		Bss3	124-150	10.5	3.6	20.1	76.2	1.66	53.0	52.0	30.1	0.20

Sand (0.02-2.0 mm); Silt (0.002-0.02mm) and Clay (<0.002mm)

formation due to wide deep surface cracks, the surface soil could have been sloughed off during rainy season and swelling pressures developed in the lower layers pushed the peds upward which leads the development of slickensides in the deeper horizons and mounds and depressions on the surface. Similar observations were made by Subbaiah and Manickam (1992); Ashok kumar and Jagadish Prasad (2010). Pressure faces were common in sub-surfaces horizon of pedon 1, 3 and 4 and slickensides were observed from 52 and 55 cm of depth, respectively and their thickness is more than 40 cm. Moderately strong to violent effervescences were observed with dilute HCl. Many calcium carbonate nodules (calcrets) were formed in lower horizons of the pedons of 1, 2, 3 and 4. The colour of $CaCO_3$ concretions vary from pale brown to light grayish white, small to bigger size (0.2 mm to 8.0 cm diameter), hard irregular outlined found in surface layers. The soft and easily separable lime nodules developed a zone of accumulation below 150 to 178 cm (Pedon 1, 3 and 4). The uniform distribution of lime concretions (and pebbles) in surface and subsurface horizon of black soils are observed. It may

Table 3	: Ratios of f	ine earth fra	ctions of pedo	ns (Particle siz	æ-analysis)				
Pedon	Location	Horizon	Depth (cm)	Sand+ Silt	Silt + Clay	Sand / Silt	Silt / Clay	Sand /(Sand + Silt)	Sand /(Silt + Clay)
	Sa dasi vpe	t di visi on							
1	Aroor	Very-Fine	Smeetitic Cale	areous Superac	t ive Isohyperth	ermic Typic H	aplusterts (Lat	it ude ° N 17 ° 37'42.34'	"Longitude [°] E 77 [°]
		53'10.33"	Altitude 539 m	ı)					
		Ар	0-25	39.9	79.4	1.02	0.33	0.50	0.25
		BA	25-52	37.4	80.5	1.03	0.30	0.51	0.24
		Bss1	52-79	34.6	81.2	1.00	0.27	0.50	0.21
		Bss2	79-115	34.0	81.5	0.98	0.27	0.49	0.21
		Bss3	115-155	36.4	81.1	0.82	0.33	0.45	0.20
		С	155+	Mixed with	calcareous mu	ram			
2	Budera	Fine Smec	titic Superactiv	e Non calcario	us Isohyperthei	mic Vertic Hap	olustepts (Latit	ude ° N 17 ° 38' 37.18"	Longitude $^{\circ}$ 77 $^{\circ}$
		50'35.00".	Altitude 585 m))					
		Ap	0-14	59.8	57.3	2.22	0.48	0.69	0.72
		Bw1	14-41	59.6	57.4	2.45	0.43	0.71	0.74
		Bwss1	41-58	61.8	55.8	2.32	0.50	0.70	0.77
		Bwss2	58-79	60.4	57.2	2.37	0.46	0.70	0.74
		BC	79-100	57.1	59.3	2.32	0.41	0.70	0.67
	Jogipet di	vision							
3	Andole	Fine Smec	titi Calcareous	Superactive Is	ohyperthermic '	Гуріс Hapluste	rts(Latitude°	17° 49'34.54" Longitu	de° 78° 05' 07.31"
		Altitude 49	92 m)						
		Ap	0-25	44.7	70.7	1.71	0.30	0.63	0.40
		BA	25-65	42.3	74.5	1.50	0.29	0.60	0.34
		Bss1	55-85	40.6	76.3	1.33	0.30	0.57	0.30
		Bss2	85-117	38.1	79.7	1.12	0.29	0.53	0.25
		Bss3	117-145	36.7	80.8	1.07	0.28	0.52	0.24
		Bss4	145-178	38.4	81.0	0.96	0.32	0.49	0.23
		С	90+	Weathered	granite- gneiss				
	Naravank	hed division							
4	Pulkurty		e Smectitic, Cal	careous Supera	active Isohypert	hermic Typic I	Haplustrets (La	titude ° 17° 56' 45.77"	Longitude° 77°
	-	•	Altitude 527 m	-			1		C
		Ар	0-27	27.8	88.9	0.59	0.25	0.37	0.12
		BA	27-55	26.3	91.5	0.45	0.25	0.31	0.09
		Bss1	55-87	24.7	93.5	0.34	0.25	0.26	0.07
		Bss2	88-124	23.8	95.2	0.22	0.26	0.18	0.05
		Bss3	124-150	23.7	96.3	0.18	0.26	0.15	0.04

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be probably due to the localized movement of the subsoil as described by Murthy *et al.* (1982). Sub surface hard pan with high bulk density was observed in sugarcane growing black soil pedon 1 and 4 where the penetration and proliferation roots are very few within 20-40 cm depth due to decades of cultivation practices of shallow ploughing and migration of finer clays to deeper layers resulted in sub soil hard pan to some extent.

Physical properties :

The clay content varied from 37.2 per cent (pedon 2) to 76.2 per cent (pedon 4). Increase of clay up to certain depth and a decrease was observed in pedon 1, 7, and 9 due to the illuviation process occurring during soil development (Table 2). Similar observations were also made by Tripathi et al. (2006). The clay content was found gradually increased in pedon 4. The increased clay content with depth was an evidence of pedogenic development as their formation and distribution is time dependent (Bhaskar et al., 2009) and also these variations could be attributed to the parent material, topography, in situ weathering and / or pedogenesis (Rudramurthy and Dasog, 2001). The silt content varied from 16.5 (pedon 3) to 19.8 (pedon 1) percent, where as in sub surface horizons ranged from 16.9 (pedon 9) to 20.1 (pedon 4) per cent. There was a gradual increase in silt content with depth in pedon 4 and no uniform trend was observed in the distribution pattern of silt content with depth in all other pedons. The silt content in most of the pedons showed an irregular trend with soil depth. It might be due to coarse nature of silt than clay, which restricts its movement with percolating water (Sharma et al., 2001).

The sand content varied from 10.3 (pedon 4) to 41.2 (pedon 2) per cent in surface horizon whereas, in subsurface horizon 3.6 (pedon 4) to 43.2 (pedon 2) per cent. A decreasing trend in sand content with depth was observed in pedon 1, 3 and 4 due to the translocation / migration of finer particles into the lower layers and surface erosion. These observations were in agreement with those of Bhaskar and Subbaiah (1995), Sarkar *et al.* (2001) and Monday *et al.* (2003). To confirm the presence or absence of lithological discontinuity among adjacent horizons in different soil pedons, the ratios of fine earth fractions were computed (Table 3). The ratios of sand / silt (0.18 (pedon 4) to 2.45 (pedon 2)), silt / clay (0.25 (pedon 4) to 0.50 (pedon 2)) and sand / (silt + clay)

(0.04 (pedon 4) to 0.77 (pedon 2)). The silt clay ratio was found to be less than 0.5 in black soils indicating the moderate weathering (Rajeshwar and Mani, 2013).

The bulk density of the soils ranged from 1.43 Mg m^{-3} (pedon 2) to 1.56 Mg m^{-3} (pedon 1) in surface horizons whereas in subsurface horizons ranged from 1.50 Mg m⁻ ³ in (pedon 2) to 1.68 Mg m⁻³ in (pedon 1). Bulk density increased with increasing depth in all pedons might be due to decrease in organic matter content, more compaction, and less aggregation (Singh and Agarwal, 2005 and Rajeshwar and Mani, 2013) except (pedon 2) showed irregular trend with soil depth. The Bss horizon of pedon 1, 3, and 4 had higher bulk density than the surface and sub-surface horizons which may be due to high clay content resulting in greater compaction in swelling clay soils (Ashok kumar and Jagadish Prasad, 2010). The higher water-holding capacity was recorded and varied from 32 per cent (pedon 2) to 53.0 per cent (pedon 3). The water holding capacity showed increasing trend with soil depth in pedon 4. The other pedons were exhibited an irregular trend with depth. In all the locations these values showed increasing trend with increasing clay content (Rudramurthy and Dasog, 2001). The pore space varied from 45.0 per cent (pedon 12) to 62.0 per cent (pedon 2). A reduction in porosity with depth was observed in pedon 1 due to soil compaction whereas reverse trend was noticed in pedon 9 and 12 (Rajeshwar and Mani, 2013).

The higher volume of expansion was varied from 23.1 per cent (pedon 1) to 30.4 per cent (pedon 3). The volume expansion was high in black soil pedons due to presence smectite type of clay minerals (Rajeshwar and Mani, 2013). The shrinkage and swelling phenomenon was exhibited by black soils, co-efficient of linear extensibility (COLE) was determined and ranged from 0.13 (pedon 3) to 0.21 (pedon 1). The studied black soils fall in the category of very high (greater than 0.09) swell-shrink class (Nayak *et al.*, 2006) might be due to increased amount of clay (Ashok kumar and Jagadish Prasad, 2010 and Rajeshwar and Mani, 2013).

Physico-chemical properties :

The pedon wise physico-chemical properties are described in table 3). The pH value of soils (pedon 1, 2, 3 and 4) found to vary from 6.5 (pedon 2) to 8.2 (pedon 1) in surface horizons (neutral to moderately alkaline) whereas in subsurface horizons ranged from 6.7 in pedon

Table	Table 4 : Physico-chemical characteristics Sugarcane growing soil pedons of the Medak district	chemical ch	n racteristi cs	Sugarcal	ne growing	soil pedor	is of the M	le dak dist	nict								
Pedon	Pedon Location	Horizon	Depth	Hq	EC 19ly	0C	B	Exchangeable Cations (c mol (p+) kg ⁻¹)	le Cations +) kg ⁻¹)		Total Ex.	BS	CEC (cmol	Free CaCO ₃	ESP	SAR	CEC/Clay
			(cm)	(07.1)	(uicip)	(g kg)	Ca	Mg	Na	K	Bases	(0%)	(p+)kg)	(%)	(0%)		railo
	Sadasivpet di vision	t division															
1	Aroor	Very-Fine	e Smectitic (alcareous)	Superactive	e Isohypert	hemic Ty	pic Haplu	sterts (Lati	itude ° N 1	7°37'42.34	"Longituc	Very-Fine Smeeticie Calcareous Superactive Isohyperthermic Typic Haplusterts (Latitude ° N 17° 37'42.34" Longitude ° E 77° 53' 10.33" Altitude 539 m)	0.33" Alt	it ude 539	(m 6	
		Ap	0-25	8.1	0.24	7.5	18.6	11.6	0.55	1.70	32.45	80.32	40.4	5.6	136	0.11	0.68
		BA	25-52	8.2	0.28	5.4	21.4	12.4	0.66	1.52	35.96	84.61	42.5	7.4	155	0.13	0.68
		Bss1	52-79	8.3	0.33	5.1	22.4	12.5	0.73	1.53	37.13	78.83	47.1	83	155	0.14	0.74
		Bss2	79-115	7.9	0.33	3.7	22.1	12.8	0.75	1.55	37.25	77.28	48.2	93	1.56	0.14	0.75
		Bss3	115-155	8.5	0.34	1.7	22.4	13.1	0.81	1.50	37.81	76.69	49.3	11.6	1.64	0.15	0.81
		С	155 +	Mixed with	ith calcareo	calcareous murram											
7	Budera	Fine Sme	ctitic Supera	ctive Non	calcarious I	sohyperth	ennic Vert	ic Haplust	epts (Latit	ude ° N 17	¹⁰ 38'37.18'	'Longitude	Fine Smeetitic Superactive Non calcarious Isohyperthernic Vertic Haplustepts (Latitude ° N 17° 38'37.18" Longitude ° 77° 50'35.00" Altitude 585 m)	00"Altitude	e 585 m)	-	
		Ap	0-14	6.5	0.21	9.9	11.4	6.0	0.14	0.76	18.30	75.00	24.4	2.8	0.57	0.04	0.63
		Bw1	14-41	6.8	0.26	4.5	12.2	6.0	0.15	0.52	18.87	66.68	28.3	3.2	0.53	0.04	0.71
		Bwss1	41-58	7.0	0.29	5.0	13.1	6.2	0.15	0.51	19.96	66.98	29.8	3.9	0.50	0.04	0.80
		Bwss2	58-79	6.7	0.31	3.5	13.2	6.4	0.20	0.55	20.35	81.08	25.1	4.7	0.80	0.05	0.64
		BC	79-100	7.0	0.34	2.4	13.3	6.5	0.21	0.55	20.56	83.92	24.5	4.7	0.86	0.05	0.58
	Jogipet division	vision															
3	Andole	Fine Sme	ctiti Calcare	ous Supera	active Isohy	perthermic	: Typic Ha	plust ert s (]	Latitude °	17° 49'34	54" Longitı	ude° 78° 05	Fine Smeetiti Calcareous Superactive Isohyperthemnic Typic Haplust etts (Latitude ° 17° 49'34.54" Longitude ° 78° 05'07.31" Altitude 492 m)	ude 492 m	(
		Ap	0-25	8.2	0.21	7.9	24.6	8.2	0.7	0.5	34.00	87.40	38.9	5.5	1.80	0.13	0.72
		BA	25-65	8.2	0.29	5.1	21.3	11.6	0.8	03	34.00	87.40	38.9	63	2.06	0.15	0.68
		Bss1	55-85	8.4	0.28	4.5	20.4	10.9	60	0.3	32.50	79.08	41.1	6.5	2.19	0.18	0.70
		Bss2	85-117	8.4	0.34	33	19.1	11.3	12	0.3	31.90	75.06	42.5	9.9	2.82	0.24	0.69
		Bss3	117-145	6.8	0.36	33	18.3	12.3	1.0	0.2	31.80	70.51	45.1	7.8	2.22	0.20	0.71
		Bss4.	145-178	9.2	0.40	2.9	18.1	8.4	13	0.3	28.10	69.69	46.3	6.6	2.81	0.28	0.75
	Narayank	Narayankhed division															
4	Pulkuty	Very-Fin	e Smectitic,	Calcareou	is Superactiv	ve Isohype	themic T	ypic Haph	ust rets (La	titude°17	° 56*45.77"	Longitude	Very-Fine Smectric, Cakareous Superactive Isohyperthermic Typic Haplust rets (Latitude ° 17° 56' 45.77" Longitude ° 77° 42' 43.82" Altitude 527 m)	2" Altitude	e 527 m)	_	
12		Ap	0-27	67	0.16	8.4	25.2	16.3	0.81	0.80	43.11	76.71	56.2	7.8	1.44	0.14	0.79
		BA	27-55	8.1	0.18	6.7	26.6	18.5	0.86	09.0	46.56	82.85	60.2	8.4	1.43	0.14	0.82
		Bss1	55-87	8.0	0.20	4.8	29.2	19.6	0.89	09.0	50.29	83.54	64.7	8.6	138	0.14	0.86
		Bss2	88-124	8.2	0.22	4.6	30.5	20.1	1.11	09.0	52.31	80.85	64.2	10.1	1.73	0.17	0.85
		Bss3	124-150	8.4	0.23	3.6	30.9	20.6	1.13	0.50	53.13	82.76	65.2	11.2	1.73	0.18	0.86

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2 to 9.2 in pedon 3 (neutral to strongly alkaline). The pH increased with depth in the pedon 3 might be due to increase in bases with depth and their complete downward leaching. The distribution was irregular in pedon 1, 2 and 4 which might be due to downward movement of bases and they get adsorbed at different layers irregularly (Rajeshwar and Mani, 2013). The EC was non saline, values found to vary from 0.16 dS m⁻¹ (pedon 4) to 0.24 dS m⁻¹ (pedon 1) in surface horizons whereas in subsurface horizons ranged from 0.18 dS m⁻¹ in (pedon 4) to 0.40 dS m⁻¹ in pedon 3. The EC gradually increased with depth might be due to the leaching of electrolytes to the lower depth and also due to foraging of nutrient ions by the vegetation in the surface layer (Renukadevi, 2003).

The organic carbon content found to vary from 6.6 g kg⁻¹ (pedon 2) to 8.4 g kg⁻¹ (pedon 4) in surface horizons whereas in subsurface horizons ranged from 1.7 g kg⁻¹ in pedon 1 to 6.7 g kg⁻¹ in pedon 4. The organic carbon content relatively higher in surface horizons than subsurface horizons in all the pedons and it decreased with depth except pedon 2. This was attributed to the addition of farmyard manure and plant residues to surface horizons which resulted in higher organic carbon content in surface horizons than that of lower horizons. These observations are in accordance with results of Rajeshwar *et al.* (2009).

The CaCO₃ content was varying from 2.8 per cent (pedon 2) to 7.8 per cent (pedon 4) in surface horizon and 3.2 per cent (pedon 2) to 11.6 per cent (pedon 1) in subsurface horizon. The content was relatively higher in deeper layers than in surface layers might be due to the downward movement of it along with percolating water (pedogenic and lithogenic) in soils of semi-arid regions (Pal *et al.*, 2000). Increase in the calcium carbonate content down the depth was attributed to the leaching of bicarbonate from upper layer during rainy season and their subsequent precipitation as carbonate in the lower layer (Maji *et al.*, 2005).

Exchangeable properties :

The cation exchange capacity was higher and varied from 24.4 c mol (p^+) kg⁻¹ (pedon 2) to 56.2 c mol (p^+) kg⁻¹ (pedon 4) in surface layers and 24.5 c mol (p^+) kg⁻¹ (pedon 2) to 65.2 c mol (p^+) kg⁻¹ (pedon 4) in sub surface layers. The high CEC of the black soils was attributed to the high clay content and smectitic clay mineralogy (Pal and Deshpande, 1987). The CEC/clay ratios were found to vary from 0.58 (pedon 7) to 0.86 (pedon 12). The CEC values are indicating that the black soils are less weathered .Higher values of CEC/clay ratio indicate the less weathered nature of the soils with weatherable primary minerals (Buol *et al.*, 1998). Soil exchange complex was dominated with Ca in all the pedons compared to other exchangeable cations and varied from 11.4 c mol (p⁺) kg⁻¹ (pedon 2) to 25.2 (pedon 4) in surface layers and 12.2 (pedon 2) to 30.9 (pedon 4) in sub surface layers. In general, exchangeable Ca content increased with depth in pedons 2 and 4 and decreased with depth in pedon 3. There was no regular pattern of distribution with depth was noticed in pedon 1.

The Mg was varied from 6.0 c mol (p^+) kg⁻¹ (pedon 2) to 16.3 c mol (p^+) kg⁻¹ (pedon 4) in surface layers and 8.4 c mol (p^+) kg⁻¹ (pedon 3) to 20.6 c mol (p^+) kg⁻¹ (pedon 4) in sub surface layers. Pedons 1, 2 and 4 showed increasing trend with soil depth. The exchangeable Na varied from varied from 0.7 c mol (p^+) kg⁻¹ (pedon 3) to 0.81 c mol (p^+) kg⁻¹ (pedon 4) in surface layers and 0.8 c mol (p^+) kg⁻¹ (pedon 3) to 1.13 c mol (p^+) kg⁻¹ (pedon 4) in sub surface layers. The exchangeable Na content increased with depth in pedons 1 and 7 and in the rest of the pedons, the depth wise distribution was irregular.

The K in soils varied from $0.5 \text{ cmol}(p^+) \text{ kg}^{-1}$ (pedon 3) to 0.80 c mol (p^+) kg⁻¹ (pedon 4) in surface layers and $0.2 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (\text{pedon } 3) \text{ to } 1.6 \text{ c} \text{ mol } (p^+) \text{ kg}^{-1} (p^+) (p^+) \text{ kg}^{-1} (p^+) (p^+)$ 1) in sub surface layers. The pedon 4 shows that the exchangeable K content decreased with depth. The remaining pedons showed inconsistent pattern with depth. The exchangeable bases in soil pedons were in order of $Ca^{+2} > Mg^{+2} > Na^{+} > K^{+}$ on the exchange complex. From the distribution of Ca⁺² and Mg⁺², it is evident that Ca⁺² shows the strongest relationship with all the species, comparing these ions (Ca⁺², Mg⁺², K⁺ and Na⁺) it was clear that Mg⁺² was present in low amount than Ca⁺² because of its higher mobility. These results are in conformity with findings of Thangasamy et al., (2005). The base saturation percentage was higher in black soil pedons ranged from 75.0 per cent (pedon 2) to 87.4 per cent (pedon 3) in surface layers and 60.69 per cent (pedon 3) to 87.40 per cent (pedon 3) in sub surface layers might be due to the dominance of smectitic type of clays and moderate to strongly alkaline reaction (Singh and Agarwal, 2005 and Gabhane et al., 2006). The SAR ranged from 0.04(pedon1) to 0.14 (pedon 4) in surface layers and 0.04 (pedon 2) to 0.28 (pedon 3) in sub surface layers .The pedons 1 and 4 found to follow an increasing trend with the increase in depth. All other pedons exhibited an irregular distribution pattern with the increase in depth.

Soil classification :

Sugar cane growing soils of Medak district were classified based on morphological, physical, physicochemical, chemical and meteorological data, according to revisions of USDA Soil Taxonomy (2010). At the highest category (order), the presence or absence of diagnostic horizons which are indications of pedogenic

processes are considered. At sub-order level, the moisture and temperature regimes were used. At lower categories (great group, family etc.) mineralogy, texture, soil chemical properties and drainage are considered. The soils of the study area were characterized and classified into two soil orders viz., Vertisols and Inceptisols.

The pedons 1, 3 and 4 were classified under Vertisols because of the following features. A layer of 25 cm or more thick, with an upper boundary with in 100 cm of the mineral soil surface; clayey texture, more than 30 per cent clay in fine earth fraction throughout the depth;

Table 5	: Avail a ble nut	trient status of	Sugarcane gr		il pedons o able macror							
Pedon	Location	Horizon	Depth	Avail	(kg ha ⁻¹)	iutrients	Available S	А	vailable mi	icron utrient	s (mg kg ⁻¹)
			(cm)	Ν	P	K	- (mg kg ⁻¹) -	Zn	Cu	Mn	Fe	В
	Sadasi vpet	di visi on										
1	Aroor	Very-Fine	Smeetitic Cale	areous Su	peractive Is	ohyperthe	micTypic Haph	usterts (Lat	it ude ° N 1	7° 37'42.34	"Longitu	de° E 77
		° 53' 10.33"	'Altitude 539 r	n)								
		Ap	0-25	289	41.8	401	44.6	0.86	2.63	7.69	8.94	1.85
		BA	25-52	210	24.1	383	31.1	0.73	2.32	7.25	7.05	1.12
		Bss1	52-79	189	19.6	281	26.2	0.54	2.60	6.65	5.52	0.56
		Bss2	79-115	168	15.2	241	16.2	0.51	1.89	5.24	4.85	0.39
		Bss3	115-155	101	10.2	222	13.3	0.36	1.48	5.44	3.27	0.42
		С	155+		Mixed	l with calca	areous murram					
2	Budera	Fine Smeet	itic Superactiv	e Non cal	carious Isol	nyperthern	nic Vertic Haplus	tept s (Lat i	tude ° N 17	° 38' 37.18'	'Longitud	e ° 77 °
		50'35.00"A	Altitude 585 m)									
		Ap	0-14	277	21.03	310	30.56	0.88	0.94	12.36	15.8	1.94
		Bw1	14-41	184	20.62	281	18.52	0.80	0.86	10.60	11.6	1.86
		Bwss1	41-58	152	14.80	267	15.47	0.42	0.66	10.25	14.4	1.51
		Bwss2	58-79	96	11.23	210	10.75	0.36	0.39	9.57	12.6	1.36
		BC	79-100	92	8.11	181	9.23	0.31	0.38	10.98	11.1	1.22
	Jogipet divi											
3	Andole	Fine Smeet	iti Calcareous	Superacti	ve Isohyper	thermic T	ypic Haplusterts (Latitude °	17°49'34.	54" Longitu	$1 de^{\circ} 78^{\circ} 0$	5'07.31"
		Altitude 49	2 m)									
		Ap	0-25	296	42.1	415	48.56	1.36	1.05	10.60	10.2	1.99
		BA	25-65	253	39.3	395	31.98	0.69	1.02	10.23	8.1	1.84
		Bss1	55-85	165	29	336	24.52	0.48	0.96	9.63	5.2	1.74
		Bss2	85-117	168	18.91	358	19.23	0.37	0.62	9.51	6.2	1.58
		Bss3	117-145	96	18.52	321	9.65	0.34	0.61	8.86	6.6	1.41
		Bss4	145-178	88	16.4	294	6.23	0.23	0.61	5.45	5.2	1.23
	Narayankh	ed di visi on										
4	Pulkurty	Very- Fine	Smectitic, Cal	careous S	uperactive	Isohyperth	ermic Typic Hap	lustrets (L	atitude °17	°56'45.77"	Lon git ud	e° 77 °
		42'43.82"	Altitude 527 m)								
		Ap	0-27	289	58.94	404	51.21	0.96	2.34	13.8	9.5	1.89
		BA	27-55	178	27.66	384	26.25	0.72	2.14	11.1	9.1	1.74
		Bss1	55-87	165	20.02	348	18.29	0.52	2.51	6.5	8.3	1.68
		Bss2	88-124	129	16.56	334	15.52	0.41	1.63	5.2	8.2	1.51
		Bss3	124-150	113	16.21	256	10.31	0.36	1.61	4.8	4.9	1.42

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gilgai micro-relief (micro-knolls and micro-ridges) on the surface; distinct intersecting slickensides in lower horizons; cracks of greater than 1 cm width which remained open and close periodically to the surface from a depth of more than 40 cm and Absence of lithic or paralithic contact, duripan, petrocalcic horizon within 50 cm from the surface. Based on these characters, the soils were grouped under order "Vertisols". The pedon 2 was classified under the order Inceptisols based on the presence of cambic subsurface horizon due to the following features. A texture of loamy very fine sand or finer; absence of rock structure in one half or more of its volume. In these pedons the subsurface horizons were recognized as cambic horizons (Thangasamy et al., 2005). As the moisture regime is Ustic, the pedons 1, 3 and 4 were classified as Usterts at sub order level. The pedons 1, 3 and 4 were further placed under the Haplusterts at great group level (Walia and Rao, 1997 and Singh et al., 1998) and Typic Haplusterts at sub group level because these pedons had deep cracks that remained open for more than 150 cumulative days most years (Surekha et al., 1997). The pedon 2 was classified as Haplustepts at great group level and Vertic Haplustepts at sub group level because of the base saturation is more than 60 per cent at a depth between 0.2 to 0.6 m from the soil surface and due to presence of cracks within 125 cm of the mineral soil surface that are 5 mm or more wide through a thickness of 30 cm or more for some time in normal years.

The pedon 1 and 4 had clay content more than 60 per cent on weighted average in fine earth fraction. Hence, it was qualified for "Very fine" particle size class. The pedon 3 had clay content less than 60 per cent on weighted average in fine earth fraction which is qualified for "fine" particle size class and showed relative dominance of smectite mineral. Hence the clay mineralogical class of these pedons was "smectitic" (Rajeshwar and Mani, 2015). The pedon 2 had showed clay content more than 35 per cent on weighted average in fine earth fraction and it was qualified for "fine" particle size class (Ramprakash and Rao,2002). The difference between mean summer and winter temperatures was less than 6° C and the mean annual soil temperature was more than 22°C. Therefore, the study area was classified as "iso- hyperthermic" temperature regime (Sehgal, 1996).

Pedogenesis :

Eastern Ghats (south) of Deccan trap had been divided into three landforms *viz.*, granite and granite - gneiss, dharawars, and cuddapahs and kurnools (Reddy *et al.*, 1996). In the very gently sloping lands and valleys, the finer factions and calcium carbonate were accumulating with weathered granitic gneiss. Hence, the parent material for the development of these black soils was weathered granite - gneiss at higher elevations and it was mixed with calcareous murram in very gently sloping lands, plains and valleys. The black soils (pedon

Table 6 : Land c	a pa bili ty clas	ssificatio	on of Suga	arcane grov	ving soil p	edons of Me	dak District bas	ed on so	il characteristi	cs			
Physiographic	Location		Topograp	ohy		Physical soil	characteristics		Pedon		l fertil: actors	2	- LCC
unit	Location	Slope	Erosion	Drainage	Texture	Sur.coarse fragments	Sub.sur.coarse fragments	Soil Depth	development	CEC	BS	OC	· Lee
Pedon 1	Aroor	II	II	IV	III	II	II	Ι	Ι	Ι	III	II	IIIwef
Pedon 2	Budera	III	II	III	II	II	II	II	Ι	Ι	Ι	III	IIIwef
Pedon 3	Andole	II	II	IV	III	II	II	Ι	Ι	Ι	Ι	II	IIIwef
Pedon 4	Pulkurty	II	II	IV	III	II	II	Ι	Ι	Ι	Ι	II	IIIwef

Table 7 : Soil-s	ite character	istics for l			sugarca	0	8				1.1		• / •	1.1	
			Clin	nate		Lanc	l form charact	eristics	Phys	sico-chemic	alchai	acteris	tics(wei	ighted av	erages)
Physiographic unit	Location	Rain fall (mm)	Max.temp (oC)	Min. temp (o C)	RH (%)	Slope (%)	Erosion	Drainage	Depth (cm)	Sur.coarse fragments (vol %)	Texture	pH(12.5)	$OC (g kg^{-1})$	CEC (Cmol (p+)/kg)	B.S (%)
Pedon 1	Aroor	855	40.0	26.2	74.0	1-3	Moderate	Poor	155	11.5	c	8.1	7.5	40.4	80.32
Pedon 2	Budera	980	40.0	26.2	74.0	1-3	Moderate	Poor	100	12.3	cl	6.5	6.6	24.4	75.00
Pedon 3	Andole	855	40.0	26.2	74.0	1-3	Moderate	Poor	178	10.4	c	8.2	7.9	38.9	87.40
Pedon 4	Pulakurty	855	40.0	26.2	74.0	1-3	Moderate	Poor	150	10.0	с	7.9	8.4	56.2	76.71

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1 and 3) were derived from weathered granite-gneiss mixed with calcareous murram whereas pedon 4 was derived from weathered basalt mixed with calcium carbonate nodules. Similar occurrence of black soils on granite-gneiss was reported earlier by Paramasivam and Gopalaswamy (1993); Subbaiah and Manickam (1992); Vijay Kumar et al. (1994). The black soils (pedon 1, 2, 3 and 4) were developed on nearly level to very gently sloping lands with slope per cent varying between 1 and 3. Many scientists in different locations also reported formation of black soils on lower elements of topography. Similar results were reported by Nagelschmidt et al. (1940) in deccan state of India; Curi and Franzmeir (1984) in central plateau of Brazil; Nagassa and Gebrekidan (2003) in Bako soils of Ethiopia and Gabhane et al. (2006) in Vidarbha region. The soils of the study area might have been formed during Archean period about 3800 million years back (Rao et al., 1995). Digar and Barde (1982) reported that it was during

Archean period, the black soils were developed during Cenozoic era, which included tertiary and quaternary period (Coulombe *et al.*, 1996).

Soil forming processes :

The black soils of sugarcane growing area (pedon 1, 3 and 4), prominent or distinct slickensides were noticed in the lower layers. Slickensides were originated due to sliding of one soil mass over the other due to swelling and expansion of clay minerals in wet season. They were seen as polished smooth surfaces in dry period when profile was opened up to the deeper layers. The pedogeneic process was nothing but argillo pedoturbation. Similar reports were earlier given by Mermut *et al.* (1996) and Maji *et al.* (2005). In the black soil locations, narrow to wide cracks were noticed revealing the shrinking nature of the clay minerals in dry period. The soil particles particularly clay which were loose on the surface, due to slight disturbance, wind and / or rain migrate to the

Table 8 : A	Actual and pot	tential so	il suitat	ility for	sugarc	ane gro	wingSo	oils of N	1e dak D	istrict							
Pedon No	Location	Max. temp ©	Min.Temp ©	RH ©	Slope (t)	Drainage (w)	Texture (s)	Depth (s)	CaCO ³ (s)	EC (n)	ESP (n)	pH (n)	BSP (f)	CEC (f)	OC (f)	Actual suitability	Potential Suitability
Pedon 1	Aroor	S1	S1	S 1	S1	S3	S3	S1	S1	S1	S1	S2	S1	S 1	S1	S2	S1
Pedon 2	Budera	S 1	S1	S1	S1	S2	S1	S 1	S2	S1	S1	S2	S 1	S1	S2	S2	S1
Pedon 3	Andole	S1	S1	S1	S1	S3	S 3	S 1	S1	S1	S1	S2	S 1	S 1	S1	S2	S1
Pedon 4	Pulakurty	S1	Sl	S1	S1	<u>S3</u>	S3	S1	SI	S1	S1	S1	S1	S 1	S1	S2	S1

Soil Suitability class: S_1 - Highly suitable; S_2 - Moderately suitable ; S_3 - Marginally suitable

 $Not \ Suitability \ class: \ N_1- \ Temporarily \ not \ suitable \ N_2 - Permanent \ ly \ not \ suitable \ AS - \ Actual \ Suitability \ PS - \ Potential \ Potential\$

Table 9 : O	Compara	ti ve e valua ti	on of productivity	of soils in the study a rea along with t	he man agement options
Soil type	Pedon	Location	Suitability	Major limitations	Management suggested
Black soils	5				
	1	Aroor	Moderately	Drainage, texture, runoff, erosion	Addition of river sand at 100 t ha ⁻¹ ; application of 100 cart
			suitable to	and CaCO3, high pH, sub surface	loads of red loam soil; summer deep ploughing; furrow
			highly suitable	hard pan	system to manage the surface drainage; raised beds should
	2	Andole	Moderately	Drainage, texture, runoff, erosion	be 1.2 m wide and 15 cm hight with two furrows of 30 cm
			suitable to	and high CaCO ₃ , high pH in	width on either side to drain out excess of water; pre
			highly suitable	subsurface horizon	monsoon sowing of green manures; application of farmy ard
	3	Pulakurty	Moderately	Drainage, texture, runoff, erosion	manures, composted coir pith or press mud at 25 t ha ⁴ per
			suitable to	and high CaCO3, high pH, sub	year and crop rotation. Follow site-specific nutrient
			highly suitable	surface hard pan	management.
	4	Budera	Moderately	Slope, medium OC and N and	Pre monsoon sowing of green manures; application of
			suitable to	Low Zn	farmy ard manures, composted coir pith or press mud at 25 t
			highly suitable		ha-1 per year and crop rotation. Follow site-specific nutrient
					management.

deeper layers along the sides of the cracks. This type of mechanical migration of inorganic particles in the profile was described as lessivage (Buol *et al.*, 1998). The 'B' horizon in the pedon 2 was exhibiting features of altered horizon and thereby resulted in structural / colour 'B' horizon (cambic horizon, a sub-surface diagnostic horizon). The colour of the soil was darker in dry and moist conditions due to release of iron oxides from weathering of rocks and minerals and their accumulation in the solum.

Land capability classification :

Land capability classification is an interpretive grouping of soils mainly based on the inherent soil characteristics, external land features and environmental factors that limits the use of land. The classification of units provide information on the physiography, colour, texture, structure of soil, type of clay mineral, consistence, permeability, depth of soil and soil reaction. Based on soil properties, soils were classified into land capability classes III with 'III swef' land capability sub-class due to the limitations of drainage, texture, erosion and soil fertility limitations Table 6). Similar interpretation was also made by Sarkar *et al.* (2002) and Rajeshwar and Mani (2013).

Soil site suitability evaluation for sugarcane :

Soil site suitability evaluation for any crops forms an essential part of every land use planning programme. The soils of the study area were evaluated for their suitability for growing sugar cane by the following criteria outlined by Naidu *et al.* (2006). Important parameters *viz.*, maximum and minimum temperature, relative humidity, slope, erosion, drainage, texture, coarse fragments, depth, soil reaction, EC, CaCO₃, organic carbon, CEC, ESP and BSP were taken into consideration for evaluating the suitability of crops (Table 7). The land suitability order and classes were assigned to the soils as per the guidelines given in the Frame Work of Land Evaluation (FAO, 1976).

The land is given a suitability rating depending on how well its properties meet the requirement of the crop. If all the properties match well with the crop requirements, the land is considered highly suitable: otherwise less suitable (moderate and marginal) and even not suitable depending upon the deviation of the land properties from the optimal growth requirement of the crops. The studied sugar cane growing soils vary in their suitability, according to the criteria for the determination of land suitability classes (Table 8).

The suitability for sugarcane cultivation in black soil pedons (1,2,3 and 4) were found moderately suitable to highly suitable with limitation of poor drainage, texture, runoff, erosion and CaCO₃, high pH, sub surface hard pan, Slope, medium OC and N and Low Zn (Table 9) with slow permeability and low hydraulic conductivity. In the black soil area, improved management practices have good potential to enhance productivity on these soils. If the improvements could be done such as addition of river sand @ 100 t ha⁻¹ and application of 100 cartloads of red loam soil; deep ploughing the field with mould board plough or disc plough during summer to enhance moderately suitable to highly suitable for the cultivation of sugarcane.

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REFERENCES

Ashok Kumar, H.P and Prasad, J. (2010). Some typical sugarcane-growing soils of Ahmadnagar district of Maharashtra: Their characterization and classification and nutritional status of soils and plants. *J. Indian Soc. Soil Sci.*, **58**(3): 257-266.

Balpande, H.S., Challa, O. and Prasad, J. (2007). Characterization and classification of grape-growing soils in Nasik district, Maharashtra. *J. Indian Soc. Soil Sci.*, **55**: 80-83.

Bhaskar, B.P. and Subbaiah, G.V. (1995). Genesis, characterization and classification of laterites and associated soils along the east coast of Andhra Pradesh. *J. Indian Soc. Soil Sci.*, **43**(1): 107-112.

Bhaskar, B.P., Baruah, U., Vadivelu, S., Raja, P. and Sarkar, D. (2009). Pedogenesis in some subaqueous soils of Brahmaputra valley of Assam. *J. Indian Soc. Soil Sci.*, **57**(3): 237-244.

Blake, G. R. and Hartze, K.H. (1986). Bulk density In Methods of Soil analysis part I(Ed A Klute). *American Society of Agronomy Incorporation* Wisconsin USA. Pp: 377-382.

Buol, S.W., Hole, F.D., Mc Cracken, R.J and Southard, R.J. (1998). *Soil geneis and classification*. (4th Ed.), Panima publishing corporation, New Delhi, pp.110.

Chapman, H.D. (1965). Cation exchange capacity. In C. A. Black et al. (ed.) Methods of soil analysis. Agronomy 9:891-901. Am. Soc. of Agron., Inc., Madison, Wis.

Chinchmalatpure, A.N., Nayak, A.K. and Rao, G.G (2005). Soil survey interpretation of salt affected black soils of Jambusar taluk of Bharuch district of Gujarat state for suggested land use planning. *Agropedol.*, **15** : 22-28.

Coulombe, C.E., Wilding, I.P. and Dixon, J.B. (1996). Over view of Vertisols. Characteristics and impacts on society. *Adv. Agron.*, **57**: 289-296.

Curi, N. and Franzmeir, D.P. (1984). Toposequence of soils from the central plateau of Brazil. *Soil Sci. Soc. American J.*, **48**: 341-346.

Digar, S. and Barde, N.K. (1982). Morphology, genesis and classification of red and lateritic soils. Review of Soil Research in India, Part - II *In*: Proceedings of the 12th International Congress of Soil Science, New Delhi, India, pp.498-507.

FAO (1976). A frame work of land evaluation. *Soils bull.*, 22. Food and Agricuture Organization of the United Nations, Rome.

Gabhane, V.V., Jadhao, V.O. and Nagdeve, M.B. (2006). Land evaluation for land use planning of a micro-watershed in Vidarbha region of Maharashtra. *J. Indian Soc. Soil Sci.*, **54**: 307-315.

Govindarajan, S.V. and Koppar, A.L. (1975). Improved method for determination of gravels in red clay soils. *J. Indian Soc. Soil Sci.*, 23: 138-140.

Jackson, M.L. (1973). *Soil chemical analysis* – Oxford IBH publishing house, Bombay, pp. 38.

Maji, A.K., Obireddy, G.P., Thayalan, S. and Walke, N.J. (2005). Characterization and classification of landforms and soils over basaltic terrain in sub humid tropics of central India. *J. Indian Soc. Soil Sci.*, **53**(2): 154-162.

Marathe, R.A., Mohanty, S. and Singh, S. (2003). Soil characterization in relation to growth and yield of Nagpur Mandarin (*Citrus reticulata* Blanco). *J. Indian Soc. Soil Sci.*, **51**(1): 70-73.

Mermut, A.R., Padmanabham, E., Eswaran, H. and Dasog, G.S. (1996). Pedogenesis. *In* Vertisols and technologies for their management developments in soil science, Ed. N Ahmad and A Mermut, Elsevier, Amsterdam, pp.32.

Monday, O., Mbila, Micheal, L., Thompson, Mbagwu, J. and David, A. (2003). Morphological and chemical properties of selected sludge amended Nigerian soil. *Soil Sci.*, **168**: 660-669.

Murthy, R.S., Bhattacharjee, J.C., Landey, R.J and Pofali, R.M. (1982). Vertisols and rice soils of tropics. 12th international congress of soil science, New Delhi, pp.3 -22.

Nagelschmidt, G., Desai, A.D. and Muir, A. (1940). The minerals in the clay fraction of a blackcotton soil and red earth from Hyderabad, Deccan state, India. *J. Indian Soc. Soil Sci.*, **30**:639-653.

Nayak, A.K., Chinchmalatpure, A.R., Rao, GG, Kandelwal, M.K. and Verma, A.K. (2006). Swell-shrink potential of Vertisols in relation to clay content and exchangeable sodium under different ionic environment. *J. Indian Soc. Soil Sci.*, **54**: 1-6.

Pal, D.K. and Deshpande, S.B. (1987). Characteristics and genesis of minerals in some benchmark Vertisols of India. *Pedol.*, **37**: 259-275.

Pal, D.K., Dasog, G.S., Vadivelu, S., Ahuja, R.L and Bhattacharyya, T. (2000). Secondary calcium carbonate in soils of arid and semi-arid rigions of India - global climate change and pedogenic carbonate. Lewis publishers, pp.149-185.

Paramasivam, P. and Gopalaswamy, A. (1993). Charactersitics and classification of some soils of lower Bhavani project command area, Tamilnadu. *Agropedol*, **3**: 105-109.

Piper, C.S. (1966). *Soil and plant analysis*. Inter Science publication. Inc. NEW YORK, U.S.A.

Rajeshwar, M. and Mani, S. (2013). Soil quality assessment in black soils of Veppanthattai Sivaganga district of Tamil Nadu. *An Asian J. Soil Sci.*, **8**(1): 108-121

Rajeshwar, M. and Mani, S. (2015). Genesis, classification and evaluation of cotton growing soils in semi arid tropics of Tamil Nadu. *An Asian J. Soil Sci.*, **10**(1): 13-22.

Rajeshwar, M., Aariff Khan, M.A. and Ramulu, V. (2009). Characterization and classification of soils of Ganapavaram pilot area of Nagarjuna Sagar Left Canal Command Area of Andhra Pradesh. *Internat. J. Tropical Agric.*, **27**: 1-7.

Ramprakash, T. and Rao, Seshagiri M. (2002). Characterization and classification of some soils in a part of Krishna district, Andhra Pradesh. *Andhra Agric. J.*, **49**: 228-236.

Renukadevi, A. (2003). Soil resource and land evaluation studies using remote sensing and GIS techniques in western Parambikulam Aliyar project area, Ph.D., Thesis, Tamil Nadu Agricultural University, COIMBATORE, T.N. (INDIA).

Richards, L.A. (1954). Diagnosis and improvement of saline and alkali soils. USDA Handbook, 60, USDA, Washington. D.C., USA.

Rudramurthy, H. V. and Dasog, G.S. (2001). Properties and genesis of associated red and black soils in north Karnataka. *J. Indian Soc. Soil Sci.*, **49**: 301-309.

Santsingh, S. (1987). Colouration and formation of black soils. *J. Agric. & Food Chem.*, **20** : 201-222. **Sarkar, D.,** Gangopadhyay, S.K. and Velayutham, M. (2001). Soil toposequence relationship and classification in lower outlier of Chhotanagpur plateau. *Agropedol.*, **11**: 29-36.

Sehgal, J.L. (1996). Pedology: Concepts and Applications. Kalyani publishers, New Delhi, pp. 176-185.

Sharma, J.P., Landey, R.J., Kalbande, A.R. and Mandal. C. (2001). Characteristics and classification of soils of Kathiawar region of Gujarat as influenced by topography. *Agropedol.*, 11:83-90.

Shyampura, R.L., Giri, J.D., Das, K. and Singh, S.K. (1994). Soil physiographic relationships on a transect in southern Rajsthan. *J. Indian Soc. Soil Sci.*, **42** : 622-625.

Singh, I.S. and Agrawal, H.P. (2005). Characterization, genesis and classification of rice soils of eastern region of Varanasi, Uttar Pradesh. *Agropedol.*, **15**: 29-38.

Soil Survey Staff. (2010). Keys to Soil Taxonomy. U.S. Dept. Agric., Natural Resources Conservation Service, Washington D.C. Oxford & IBH Publishing Co., NEW DELHI, INDIA.

Srinivasan, T.R., Balu Y.P and Tamhane, R.V. (1969). Placement of black soils of India in the comprehensive soil classification

system-7th approximation. J. Indian Soc. Soil Sci., 17: 323-332.

Subbiah, G.V. and Manickam, T.S. (1992). Genesis and morphology of Vertisols developed on different parent materials. *J. Indian Soc. Soil Sci.*, **40**: 150-155.

Surekha, K., Subbarao, I.V., Rao, A.P and Shantaram, M.V. (1997). Characterization of some Vertisols of Andhra Pradesh. *J. Indian Soc. Soil Sci.*, **45:** 338-343.

Thangasamy, A., Naidu, M.V.S, Ramavatharam, N. and Raghava Reddy, C. (2005). Characterization, classification and evaluation of soil resources in Sivagiri micro watershed of Chittoor district in Andra Pradesh for sustainable land use planning. *J. Indian Soc. Soil Sci.*, **53**(1): 11-21.

Vijaya Kumar, T., Reddy, M.S. and Gopalakrishna, V. (1994). Characteristics and classification of soils of northern Telangana zone of Andhra Pradesh. *Agropedol.*, **4**: 31-43.

Walkely, A.J. and Black, C.A. (1934). An estimation of the digestion method for determining soil organic matter and a proposed modification of the chromic acid titration method. *Soil Sci.*, **37**: 29-38.

